# **Short Communication:**

# Study of Environmental Isotopes and Hydrochemical Characteristics of Groundwater from Springs at Archaeological Sites in Dompu Regency, West Nusa Tenggara, Indonesia

# Satrio Satrio<sup>1\*</sup>, I Nyoman Rema<sup>2</sup>, Sonny Christophorus Wibisono<sup>2</sup>, Luh Suwita Utami<sup>2</sup>, Nyoman Arisanti<sup>3</sup>, I Gusti Ngurah Jayanti<sup>3</sup>, and I Wayan Rupa<sup>4</sup>

<sup>1</sup>Research Center for Radiation Process Technology, Research Organization for Nuclear Energy, National Research and Innovation Agency (BRIN), Jakarta 12440, Indonesia

<sup>2</sup>Research Center for Prehistoric and History Archaeology, Research Organization for Archaeology, Language and Literature, National Research and Innovation Agency (BRIN), Jakarta 12510, Indonesia

<sup>3</sup>*Research Center for Environmental Archaeology, Maritime Archaeology and Cultural Sustainability, Research Organization for Archaeology, Language and Literature, National Research and Innovation Agency (BRIN), Denpasar 80223, Indonesia* 

<sup>4</sup>Treasures of Religion and Civilization, Research Organization for Archaeology, Language and Literature, National Research and Innovation Agency (BRIN), Denpasar 80223, Indonesia

#### \* Corresponding author:

email: satr002@brin.go.id Received: April 10, 2023 Accepted: October 16, 2023

DOI: 10.22146/ijc.83792

**Abstract:** The existence of groundwater sources in several springs at archaeological sites in Dompu Regency, West Nusa Tenggara, Indonesia, has been widely used by the surrounding community for various needs. However, from a number of the springs, there are springs whose water discharge has decreased. Meanwhile, from a number of existing springs, there is one spring whose groundwater is used every day even though it tastes a bit brackish. For this reason, it is important to conduct a groundwater study in the area with the aim of knowing the characteristics, preliminary identification of recharge areas and quality of groundwater in the study area through an environmental isotope and hydrochemical. The study was conducted by taking a number of groundwater samples from several archaeological sites in Dompu Regency. The results of environmental isotope and hydrochemical analysis show that there are 2 springs (2 archaeological sites), namely the Riwo and Ncona springs, because these two areas are part of the recharge area, which must be preserved by not clearing forest land. Meanwhile, for the quality of groundwater, of the 5 springs located at the archaeological sites, only the Hodo spring is of "poor quality" with the Na–Cl water type; it is unfit for drinking water.

*Keywords:* Dompu groundwater; groundwater characteristics; isotope, hydrochemical; Dompu archaeological sites; water quality

#### INTRODUCTION

In several locations of archaeological sites in Dompu Regency, West Nusa Tenggara, Indonesia, there are springs that are still used by the surrounding community. This use has been going on for a long time, since the Ncuhi period until now. Ncuhi is a tribal community group, according to several sources of information from the Dompu community, which existed before the time of the Dompu kingdom. During the Ncuhi era, although they were still in small groups, the people routinely interacted with the surrounding nature, choosing high places such as hills and mountains as a place to live, which were located close to rivers or springs [1-3]. Meanwhile, in the area around the lowlands near the coast, there is a settlement called Ncuhi Tonda, which also has several springs. The location of the springs is part of an archaeological site, traces of the Dompu kingdom in the past. Until now, some of the springs have been widely used by the local community for various daily needs such as for drinking water, washing and ritual places. However, not all groundwater from springs located at the archaeological sites can be used to meet daily needs, especially for drinking water, because there are springs whose water is unfit for consumption. On the other hand, over time, followed by the development of population, settlements, and the large number of clearings of forest land for plantations, this has the effect of reducing groundwater recharge areas, which greatly affects the decrease in groundwater discharge in some springs at the sites. Also, this causes a reduction in the availability of water that supplies agricultural irrigation. For these reasons, it is very important to conduct a groundwater study in the area with the aim of knowing the characteristics of groundwater (such as meteoric water, evaporation water, water-rock interaction, water type, and hydrochemical control of groundwater), preliminary identification of recharge areas and also to determine the quality of groundwater, especially groundwater that is widely used by the community, both for drinking water and for other daily needs. This study was conducted using environmental isotope (<sup>2</sup>H, <sup>18</sup>O, <sup>14</sup>C) and hydrochemical approaches (Na<sup>+</sup>, K<sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, Cl<sup>-</sup>, SO<sub>4</sub><sup>2-</sup>, HCO<sub>3</sub><sup>-</sup>, NO<sub>3</sub><sup>-</sup>).

#### EXPERIMENTAL SECTION

#### Materials

The materials used in this study were BaCl<sub>2</sub>, NaOH, FeSO<sub>4</sub>·7H<sub>2</sub>O, ethanol (Merck, Germany), polyacrilamide (BDH Chemicals Ltd, England), Carbosorb-E, Permafluor-E (Perkin Elmer, USA), and liquid nitrogen.

#### Instrumentation

The instrumentations used in this study were liquid water isotope analyzer DLT-100 (LGR, USA), HTS PAL autosampler (LEAP Technologies, USA), liquid scintillation analyzer 2910TR (Perkin Elmer, USA), ion chromatography 883 Basic IC plus (Metrohm, Switzerland), and automatic potentiometric titrator AT-710 (KEM, Japan).

#### **Study Area**

The study area is located in several sub-districts in Dompu Regency, West Nusa Tenggara, Sumbawa Island. Geographically, the study area is located at coordinates between 577122.37-664851.09 mE and between 8995593.97-9104580.58 mS. Archaeologically, the study area is a trace of ancient settlements and is generally located close to water sources. Study activities were carried out in the Districts of Woja, Pajo and Pekat. In Woja District, namely in Riwo Village, there is an important site located in the Tonda hills which has a spring at the foot of the hill, namely the Riwo spring, which is not too far from Cempi Bay. Not too far from the Riwo Spring, there is the Ncona Spring located in Temba Lae Village, Pajo District, which is also not too far from Cempi Bay. Several locations that were considered important in the past include the beach, which has springs, namely in Pekat District, south of Mount Tambora, which has three springs, namely the Hodo, Rao and Wau springs. A map of the location for taking water samples from several springs in Dompu Regency can be seen in Fig. 1.

#### Procedure

#### Water sampling

Sampling for environmental isotope analysis of <sup>18</sup>O and <sup>2</sup>H was carried out by taking 30 mL of groundwater samples into an airtight bottle. Air bubbles in bottles should be avoided to prevent evaporation [4]. For environmental radioactive isotope analysis, <sup>14</sup>C was taken in the form of precipitated BaCO<sub>3</sub> carbonate, which was put into a 1 L plastic bottle. Meanwhile, for hydrochemical analysis, about 300 mL was taken, which was put into a plastic bottle.

#### Analysis of environmental isotopes (<sup>2</sup>H, <sup>18</sup>O, <sup>14</sup>C)

Analysis of stable isotopes was done by laser spectroscopic method, i.e., using Los Gatos research (LGR) DLT-100 liquid water isotope analyzer. The composition of isotopes is expressed as relative ratio ( $\delta$ ) against standard mean oceanic water (SMOW) as Eq. (1) [5];

$$\delta = \frac{\text{Rsample} - \text{RSMOW}}{\text{RSMOW}} \times 1000\%$$
(1)

where  $\delta{:}\,\delta^2 H$  or  $\delta^{18} O$  (in unit ‰) and R :  $^2 H/^1 H$  or  $^{18} O/^{16} O.$ 



Fig 1. Map of the location of groundwater sampling in the study area

The <sup>14</sup>C isotope analysis was carried out using the  $CO_2$  Absorption method on a series of absorption line devices. This method involves a solution of Carbosorb-E/Permafluor-E to absorb  $CO_2$  into a solution of carbamate [6]. Meanwhile, the counting process for the <sup>14</sup>C isotope in the carbamate solution was carried out using a liquid scintillation analyzer (LSA).

#### Analysis of hydrochemical

Analysis of hydrochemicals was done by the following methods: for  $HCO_3^-$  using autotitrator, while for  $Cl^-$ ,  $SO_4^{2-}$ ,  $F^-$ ,  $NO_3^-$ ,  $Na^+$ ,  $K^+$ ,  $Ca^{2+}$ ,  $Mg^{2+}$ , using ion chromatography.

#### RESULTS AND DISCUSSION

#### **Environmental Isotope Characteristic**

Table 1 shows the analysis results of stable environmental isotopes (<sup>18</sup>O, <sup>2</sup>H) and environmental radioactive isotope <sup>14</sup>C from groundwater samples taken from springs at archaeological sites in the study area. From the results of the <sup>14</sup>C isotope analysis of groundwater dating from water samples taken from several spring locations in the study area, it can be seen that the age of the groundwater is 320 years BP, and the oldest is 3860 years BP (BP = before present = before 1950, the modern reference international standard for measuring the isotope <sup>14</sup>C, where 100 percent Modern Carbon = 0.95 activity of oxalic acid standard). However, the results of this groundwater age analysis must be corrected according to the geological conditions of the study area [7]. According to Geyh, groundwater discharge somewhere does not reflect 100 percent Modern Carbon (pMC), during infiltration [8], this is due to the influence of the dissolution of carbonate from rock along its path. Dissolved carbonates from rocks tend to give a much older age than they should [9]. For this reason, it is necessary to correct the age of groundwater according to the dominant rock type in the study area so that the corrected age is obtained.

The study area is a volcanic area, so to correct the corrected age results, the initial activity between 90-100 pMC is taken. If the initial activity of 95 pMC is taken, then the corrected age data is obtained, as can be seen in Table 1. It can be seen that groundwater from Riwo and Ncona springs have groundwater age of 115 years BP and Modern, respectively, classified as young groundwater age, indicating that the area around the two springs is part of the recharge area of groundwater so that the area should be reforested by planting trees that are able to bind water during the rainy season and most of it can be deposited in groundwater aquifers [10]. The Hodo spring has a relatively old water age of 3550 years BP or has a longer storage period in the aquifer than other springs [11]. Meanwhile, the Rau and Wau springs have a relatively younger age than the Hodo springs,

Springs location	Coordinate	Elevation (m asl)	δ <sup>18</sup> Ο (‰)	δ²H (‰)	<sup>14</sup> C activity (PMC)	<sup>14</sup> C Age (uncorrected, yr BP)	<sup>14</sup> C Age (empirical model with A <sub>0</sub> =95 pMC, yr BP)
Hodo spring	619058.83 mE, 9065746.13 mS	1	-5.17	-33.2	61.84	3860	3550
Rau spring	598116.94 mE, 9071183.80 mS	1	-5.61	-32.9	79.97	1795	1425
Wau spring	585285.98 mE, 9082320.33 mS	1	-5.77	-34.0	81.16	1675	1300
Riwo spring	646211.84 mE, 9041899.79 mS	163	-4.82	-28.0	93.71	520	115
Ncona spring	664987.79 mE, 9046194.42 mS	139	-4.05	-22.3	96.08	320	Modern

Table 1. The results of isotopes of <sup>18</sup>O, <sup>2</sup>H and <sup>14</sup>C analysis of the springs in the study area

namely 1425 years BP and 1300 years BP, respectively; it is also very important to conserve the recharge area. The Hodo, Rau, and Wau springs probably originate from the recharge area on the slopes of Mount Tambora.

Based on the analysis of stable environmental isotopes of <sup>18</sup>O and <sup>2</sup>H, it can be seen that the variation of isotope values for  $\delta^{18}$ O and  $\delta^{2}$ H in the study area varies from -5.61 to -4.05‰ for  $\delta^{18}$ O and between -33.2 to -22.3‰ for  $\delta^{2}$ H. Furthermore, the values of the isotopes  $\delta^{18}$ O and  $\delta^{2}$ H are plotted in the form of a graph of the relationship between  $\delta^{2}$ H vs  $\delta^{18}$ O, as shown in Fig. 2. From the graph, it can be seen that Wau and Rau's springs have relatively the same environmental isotope values of  $\delta^{18}$ O and  $\delta^{2}$ H, this indicates that the water from the two springs originates from the relatively similar recharge elevation on the slopes of Mount Tambora with the age of

the groundwater is also relatively similar as described above. Also, Fig. 2 shows that the Hodo spring probably originates from the same infiltration elevation as the Wau and Rau springs, but when it enters the groundwater aquifer system, it passes through the deeper aquifer layers and dissolves silicate minerals such as plagioclase feldspar {(Ca,Na)AlSi<sub>3</sub>O<sub>8</sub>} at the depth of the aquifer so that the groundwater of the Hodo spring shows the phenomenon of oxygen shift due to the interaction between oxygen from water and oxygen from silicate minerals [12-13]. The Hodo spring groundwater flow through this deeper aquifer layer is supported by its <sup>14</sup>C isotope data, which results in a relatively older groundwater age compared to other springs as described above. Furthermore, for Riwo Spring, which is at an elevation of 163 m above sea level and the groundwater



Fig 2. Isotope graph of  $\delta^2$ H vs  $\delta^{18}$ O groundwater from springs in the study area

age is around 115 years BP, it indicates that the recharge area is at an elevation above it. Because the age of the water is relatively young, the area at an elevation above the spring is a recharge area that must be preserved. Likewise, with the Ncona spring, the groundwater is of the modern age, and the isotopic value ( $\delta^{18}$ O,  $\delta^{2}$ H) is more enriched than the previous 4 springs; the Ncona spring area is part of a groundwater recharge area that really needs to be preserved for trees that are in the vicinity and no new land is cleared for plantations or settlement.

#### Hydrochemical Characteristic

#### Hydrochemical composition and water type

Table 2 shows the results of hydrochemical analysis of groundwater from springs at archaeological sites in the study area with an average ionic balance of 2.69%. In general, the groundwater of Hodo Spring has a higher cation-anion composition than other springs. This is in line with the results of isotopic analysis of <sup>18</sup>O, <sup>2</sup>H and <sup>14</sup>C, which indicated that the Hodo Spring originated from deep aquifer flows which were characterized by high concentrations of Na<sup>+</sup> reaching 300.58 mg/L and Cl<sup>-</sup> at 470.45 mg/L. In addition, the high sulfate (SO<sub>4</sub><sup>2-</sup>) in the Hodo Spring is probably caused by deposits of sulfate minerals (such as gypsum, CaSO<sub>4</sub>·2H<sub>2</sub>O) [14], which are

deposits of the remnants of the past eruption of Mount Tambora dissolved in groundwater flows. The concentration of Na<sup>+</sup> and Cl<sup>-</sup> ions from Hodo Spring also exceeds the threshold for drinking water health requirements based on the Minister of Health of the Republic of Indonesia No. 492 in 2010, and the WHO standard in 2011 are 200 and 250 mg/L, respectively. The hydrochemical compositions of other springs, namely Rau, Wau, Riwo, and Ncona, have ion concentrations below the threshold for drinking water health requirements, so they are safe for drinking water and for other daily needs.

Based on the Piper diagram (Fig. 3), it is shown that groundwater from the Hodo spring has the water type of Na–Cl, which is the character of salt water or brackish water. Meanwhile, for Ncona and Riwo springs, the water type is Ca–Mg–HCO<sub>3</sub>, indicating the dominant hydrochemical composition originated from dissolving minerals such as dolomite. The water type of Ca–Mg– HCO<sub>3</sub> from the two springs also indicates that the surrounding area is a groundwater recharge area for groundwater in the discharge area at the elevation below. This is in line with the results of isotope analysis, especially the <sup>14</sup>C isotope, which obtained a Modern age and close to Modern age. Furthermore, Rau spring has the



Fig 3. Piper diagram of groundwater from springs in the study area

Spring			С	ation-a	nion (mg	;/L)			%	TDC	Total	пU
location	Na <sup>+</sup>	$K^+$	Ca <sup>2+</sup>	$Mg^{2+}$	Cl-	$\text{HCO}_3^-$	$\mathrm{SO_4}^{2-}$	$NO_3^-$	balance	103	hardness	PII
Hodo	300.58	50.02	27.59	30.34	470.45	211.12	160.61	1.36	4.87	1254	193.8	7.31
Rau	24.44	15.22	14.57	16.81	2.72	204.54	16.23	2.36	3.29	298	105.6	6.98
Wau	34.14	9.51	7.05	9.16	18.09	142.00	12.55	0.88	4.69	235	55.3	7.90
Riwo	22.95	4.37	27.11	29.82	15.58	263.43	7.62	0.35	0.04	371	190.5	7.14
Ncona	12.58	2.09	26.92	26.92	3.02	229.40	17.27	0.01	0.54	319	178.1	6.28

Table 2. The results of hydrochemical analysis of groundwater in the study area

Table 3. Hydrochemical characteristics of spring groundwater in the study area

Spring	Cation-anion	Water type	Groundwater
location	concentration		status
Hodo	Na, Cl and SO <sub>4</sub> exceed the threshold	Na-Cl	Brackish
Rau	Normal	Ca-HCO <sub>3</sub>	Freshwater
Wau	Normal	Mixed type	Freshwater
Riwo	Normal	Ca-Mg-HCO <sub>3</sub>	Freshwater
Ncona	Normal	Ca-Mg-HCO <sub>3</sub>	Freshwater

water type of Ca–HCO<sub>3</sub>, which is the dominant hydrochemical composition originating from the dissolution of calcite minerals. Meanwhile, Wau spring has mixed Ca–Na–HCO<sub>3</sub> water type with dominant Na<sup>+</sup> and HCO<sub>3</sub><sup>-</sup> ions, indicating an exchange of Ca<sup>2+</sup> or Mg<sup>2+</sup> ions from groundwater replaced by Na<sup>+</sup> from silicate minerals through a freshening process so that the concentration of Na<sup>+</sup> ions in groundwater increases. The different water types between Hodo spring (Na–Cl) and other springs indicate differences in genesis and differences in the aquifer through which groundwater passes, as described in the discussion above [15]. Table 3 shows the hydrochemical characteristics based on the cation-anion concentration, water type and groundwater status of the springs in the study area.

# Hydrochemical control of groundwater

Based on the Gibbs diagram, in general, there are three types of processes that control the chemical mechanism of groundwater, namely precipitation or rainfall, evaporation crystallization, and rock-water weathering in a study area [16-17]. Gibbs made a simple but effective diagram using TDS vs Na<sup>+</sup>/(Na<sup>+</sup> + Ca<sup>2+</sup>) and TDS vs Cl<sup>-</sup>/(Cl<sup>-</sup> + HCO<sub>3</sub><sup>-</sup>) ratios to identify factors influencing the hydrochemical of groundwater [18-19]. Fig. 4(a) and 4(b) show the hydrochemical processes in water samples from springs in the study area. It can be seen that the hydrochemical of Hodo spring is controlled by the dissolution process or rock-water weathering dominance. While other springs, especially Riwo and Ncona springs, tend to be controlled by rainwater or depend on relatively high rainwater, during long dry seasons, not much groundwater reserves are deposited or stored properly in the aquifer. This condition is probably due to reduced recharge areas around the two springs due to reduced forest land. The following Table 3 shows the hydrochemical characteristics based on the cation-anion concentration, water type and groundwater status of the springs in the study area.

# Groundwater quality

Another characteristic of the five springs is the water quality index (WQI), which is calculated based on the hydrochemical data above. Table 4 shows the hydrochemical parameters used to calculate WQI of the groundwater in the study area.

As a reference in determining the WQI, the 2011 WHO international standard on the 2011 WHO guidelines for drinking water quality was used [20-21]. Based on the results of the WQI calculation, as can be seen in Table 5, Hodo Spring it has a value of 140.24 with the status of "poor water" or water that is not suitable for drinking water, while groundwater from other springs has the status of "good water" and "excellent water".



Fig 4. Hydrochemical control of groundwater from springs in the study area

Table 4. The hydrochemical parameters for calculating WQI of groundwater from springs in the study area

Parameter	Weight	Relative weight	Maximum concentration
	$(W_i)$	(W <sub>i</sub> )	(WHO Standard, 2011), C <sub>i</sub>
Na <sup>+</sup>	4	0.129	200 mg/L
$K^+$	3	0.097	12 mg/L
Ca <sup>2+</sup>	4	0.129	100 mg/L
$Mg^{2+}$	4	0.129	50 mg/L
Cl-	3	0.097	250 mg/L
$HCO_3^-$	2	0.065	125 mg/L
$\mathrm{SO_4}^{2-}$	1	0.032	250 mg/L
$NO_3^-$	1	0.032	50 mg/L
Hardness	2	0.065	500 mg/L
pН	4	0.129	6.5-8.5
TDS	3	0.097	600 mg/L

**Table 5.** Calculation results of groundwater quality from springs in the study area

<u> </u>		
Spring location	WQI	Groundwater quality
Hodo	140.24	Poor water
Wau	41.07	Excellent water
Rau	51.21	Good water
Riwo	53.86	Good water
Ncona	45.77	Excellent water

## CONCLUSION

Isotope and hydrochemical characteristics of several springs at archaeological sites in Dompu Regency, i.e., Riwo and Ncona springs with Ca-Mg-HCO<sub>3</sub> water type have a groundwater age of 115 years BP and Modern,

respectively, classified as young groundwater age. The area of these two springs is part of the recharge area, which is controlled by the presence of rainwater, so this area really needs reforestation, and no new land clearing is carried out in the existing forest. Hodo spring is typical of brackish water with Na-Cl water type, the age is relatively old compared to other springs and flows Meanwhile, through deep aquifers. the high concentration of SO4<sup>2-</sup> ion but has a relatively neutral pH in the Hodo spring is thought to come from sulfate mineral deposits (such as gypsum) which are dissolved in groundwater. The sulfate mineral deposit itself is estimated to be a product of the volcanic activity of Mount Tambora in the past; The results of the

calculation of the water quality index, except for the Hodo spring, other springs are still classified as freshwater which fulfills health requirements as groundwater suitable for drinking water.

### ACKNOWLEDGMENTS

This research was funded by the Regional Regency for Archaeological Research in Bali Province, Ministry of Education, Culture, Research & Technology, DIPA-023.11.2.690387/2021. We would like to deliver our grateful thanks to those who have helped us: Syafrudin from the Dompu Regency Environmental Service, Nurhaidah as Dompu Historian, Abubakar and Rian Eko Muslimin from the Department of Public Works and Spatial Planning, Tarmizi from the Tourism and Culture Office of Dompu Regency, who helped us during the observation.

## AUTHOR CONTRIBUTIONS

I Nyoman Rema designed and directed the study. Sonny Christophorus Wibisono and Luh Suwita Utami provide a description of the Dompu civilization and the Dompu archaeological site. Nyoman Arisanti, I Gusti Ngurah Jayanti and I Wayan Rupa were involved in sampling in the study area. Satrio carried out data processing, data interpretation and wrote articles.

#### REFERENCES

- [1] Vansina, J., 1985, *Oral Tradition as History*, University Wisconsin Press, Wisconsin, US.
- [2] Hagerdal, H., 2017, Held's History of Sumbawa, Amsterdam University Press B.V., Amsterdam, Netherland.
- [3] Rema, I.N., Juliawati, N.P., and Prihatmoko, H.,
   2018, Doro Bata site in Dompu, Nusa Tenggara Barat: Study form, space, and time, *Kapata Arkeologi*, 14 (1), 79–88.
- [4] Meng, Y., Liu, G., and Li, M., 2015, Tracing the sources and processes of groundwater in an Alpine glacierized region in Southwest China: Evidence from environmental isotopes, *Water*, 7 (6), 2673–2690.
- [5] Satrio, S., Prasetio, R., Alam, B.Y.C.S.S.S., Hadian, M.S.D., Hendarmawan, H., 2021, Comparison of isotope and hydrochemical characteristics of springs

in Sembalun – Rinjani Area, East Lombok, West Nusa Tenggara, Indonesia before and after the earthquake events in 2018, *Bull. Geol. Soc. Malays.*, 71, 215–226.

- [6] Rema, I.N., Satrio, S., and Arisanti, N., 2022, The Utilization of the Dorobata Terrace, Dompu Regency, West Nusa Tenggara, Proceedings of the 9<sup>th</sup> Asbam International Conference (Archeology, History, & Culture in The Nature of Malay) (ASBAM 2021), Advances in Social Science, Education and Humanities Research, Atlantis Press, 660, 70–77.
- [7] Irvine, D.J., Wood, C., Cartwright, I., and Oliver, T., 2021, Depth to water table correction for initial carbon-14 activities in groundwater mean residence time estimation, *Hydrol. Earth Syst. Sci.*, 276, 5415– 5424.
- [8] Mook, W.G., 2000, Environmental Isotopes in The Hydrological Cycle: Principles and Applications, International Atomic Energy Agency and United Nations Educational, Scientific and Cultural Organization, Paris, France.
- [9] Hershey, R.L., Fereday, W., and Thomas, J.M., 2016, Dissolved Organic Carbon <sup>14</sup>C in Southern Nevada Groundwater and Implications for Groundwater Travel Times, Technical Report, Nevada Field Office, National Nuclear Security Administration, U.S. Department of Energy, Las Vegas, Nevada, Publication No. 45268.
- [10] Smerdon, B.D., Redding, T.E., and Beckers, J., 2009, An overview of the effects of forest management on groundwater hydrology, *J. Ecosyst. Manage.*, 10, 22– 44.
- [11] IAEA, 2013, Isotope Methods for Dating Old Groundwater, Isotopes in Environmental and Health Studies, International Atomic Energy Agency, Vienna, Austria.
- [12] Boschetti, T., Cifuentes, J., Iacumin, P., and Selmo,
  E., 2019, Local meteoric water line of Northern Chile (18° S-30° S): An application of error-invariables regression to the oxygen and hydrogen stable isotope ratio of precipitation, *Water*, 11 (4), 791.

- [13] Torbehbar, A.K., and Sattari, S.M., 2015, Geochemistry and Isotope Study of Discharged Geothermal Fluids, NW Sabalan Geothermal Field, NW Iran, *Proceedings World Geothermal Congress* 2015, 19-25 April 2015, Melbourne, Australia.
- [14] Luo, W., Gao, X., and Zhang, X., 2018, Geochemical processes controlling the groundwater chemistry and fluoride contamination in the Yuncheng Basin, China–An area with complex hydrogeochemical conditions, *PLoS One*, 13 (7), e0199082.
- [15] Lu, Y., Tang, C., Chen, J., and Chen, J., 2015, Groundwater recharge and hydrogeochemical evolution in Leizhou Peninsula, China, J. Chem., 2015, 427579.
- [16] Dedzo, M.G., Tsozué, D., Mimba, M.E., Teddy, F., Nembungwe, R.M., and Linida, S., 2017, Importance of rocks and their weathering products on groundwater quality in Central-East Cameroon, *Hydrology*, 4 (2), 23.
- [17] Lyu, M., Pang, Z., Yin, L., Zhang, J., Huang, T., Yang, S., Li, Z., Wang, X., and Gulbostan, T., 2019, The control of groundwater flow systems and geochemical processes on groundwater chemistry: A case study in

Wushenzhao basin, NW China, Water, 11 (4), 790.

- [18] Zhang, B., Zhao, D., Zhou, P., Qu, S., Liao, F., and Wang, G., 2020, Hydrochemical characteristics of groundwater and dominant water-rock interactions in the Delingha area, Qaidam basin, northwest China, *Water*, 12 (3), 836.
- [19] Wu, C., Wu, X., Qian, C., Zhu, G., 2018, Hydrogeochemistry and groundwater quality assessment of high fluoride levels in the Yanchi endorheic region, northwest China, *Appl. Geochem.*, 98, 404–417.
- [20] Badr, E.A., and Al-Naeem, A.A., 2021, Assessment of drinking water purification plant efficiency in Al-Hassa, eastern region of Saudi Arabia, *Sustainability*, 13 (11), 6122.
- [21] Reyes-Toscano, Claudia A., Alfaro-Cuevas-Villanueva, R., Cortés-Martínez, R., Morton-Bermea, O., Hernández-Álvarez, E., Buenrostro-Delgado, O., and Ávila-Olivera, J.A., 2020, Hydrogeochemical characteristics and assessment of drinking water quality in the urban area of drinking water quality in the urban area of Zamora, Mexico, Water, 12 (2), 556.